CALC-ALKALINE VOLCANIC ROCKS IN THE PRE-UPPER JURASSIC BASEMENT OF SOUTH-CENTRAL MEXICO—A LATE PALEOZOIC-EARLY MESOZOIC CONVERGENT PLATE MARGIN?

INTRODUCTION

The limited extension of geological map coverage, the metamorphism of varying degrees that affected the rocks and their almost complete lack of fossils, make the basement beneath the marine fossiliferous Upper Jurassic and/or Lower Cretaceous sedimentary and volcanic cover rocks to be an unsolved geological problem in south-central Mexico, a region comprising about 20,000 km² (Figure 1). Radiometric ages (Pb- α , K-Ar, Rb-Sr, U-Pb) obtained from these rocks are apparently contradictory or not conclusive (de Cserna et al., 1962; de Cserna, Fries, Rincón-Orta, Westley, Solorio-Munguía and Schmitter-Villada, 1974; de Cserna, Fries, Rincón-Orta, Solorio-Munguía and Schmitter-Villada, 1974; Guerrero-García et al., 1978; Urrutia-Fucugauchi and Linares, 1981; López-Infanzón and Grajales-Nishimura, 1984; Morán-Zenteno et al., 1990; Morán-Zenteno, 1992). The various paleogeographic and paleotectonic reconstructions advanced so far are merely hypothetical models (de Cserna, 1958, 1976; Ortega-Gutiérrez, 1981; Campa-Uranga and Coney, 1983; Sedlock et al., 1993). One important question regarding this basement refers to the nature of the plate-tectonic scenario in which these rocks were formed; accordingly, the authors have chosen to examine the chemical character of the volcanic rocks contained in the younger greenschist-facies package of the pre-Upper Jurassic and/or Lower Cretaceous basement in south-central Mexico.

GEOLOGIC SETTING

The pre-Upper Jurassic and/or Lower Cretaceous basement in south-central Mexico, roughly between meridians 99° and 100°30'W (Figure 1), consists of an older package of amphibolite-facies metasedimentary, metavolcanic, and metaplutonic complex (the Xolapa Complex; de Cserna, 1965). This complex is covered by a younger greenschist-facies package, consisting of metasedimentary and metavolcanic units (the Taxco Schist [Fries, 1960], the Ixcuinatoyac Formation [Klesse, 1968] and the Ayotusco Formation [Díaz-García, 1980]), a metavolcanic and to a lesser extent metavolcaniclastic unit (the Taxco Viejo Greenstone [Fries, 1960; de Cserna and Fries, 1981]), and a quartzitic continental redbed with Zoltan de Cserna,* Jerjes Pantoja-Alor,* Mariano Elías-Herrera,* and José Luis Sánchez-Zavala*

associated volcanic rocks unit (the Chapolapa Formation [de Cserna, 1965]). This younger greenschist-facies basement package contains several unconformities and is beveled by an erosion surface over which the Upper Jurassic and/or Lower Cretaceous marine cover was deposited. Such a cover consists chiefly of limestone, in the east, and grades westward into a carbonate-volcaniclastic-volcanic sequence that, at places, shows the effects of a low-grade (prehnite-pumpellyite facies) metamorphism (Sánchez-Zavala and Elías-Herrera, 1991).

For this inquiry, six localities in south-central Mexico were examined, which were mapped previously by the present authors (Figure 1), and collected representative rock samples from the volcanic intervals of the various lithostratigraphic units of the younger greenschist-facies basement package.

SAMPLED STRATIGRAPHIC UNITS

The so far known oldest lithostratigraphic unit of the younger greenschist-facies basement package is the Taxco Schist or its correlative unit, the Ixcuinatoyac Formation. Volcanic intervals in the Taxco Schist were sampled at two localities. The first of these is at an excellent outcrop along the Tejupilco-San Lucas del Maíz dirt road in southwestern State of Mexico (Figure 1, locality A); here, the analysed sample (Table 1, sample A) comes from a well-foliated andesite metalava (de Cserna, 1982; Elías-Herrera, 1989). The second sampling locality is about 160 km south-southeast of the previous one. It is at the left abutment of the La Venta dam on the river Papagayo (Figure 1, locality F), about 6 km southwest of Tierra Colorada in south-central State of Guerrero, and about 7 km east of the type locality of the Xolapa Complex (de Cserna, 1965). The analysed sample (Table 1, sample F) is a foliated volcaniclastic rock of dacite composition. It should be noted that the original reconnaissance mapping of this area carried out almost 40 years ago shows these rocks as Chapolapa Formation (de Cserna, 1965); however, based on our present knowledge of the regional distribution of the Taxco Schist, we consider the La Venta dam locality to be the Taxco Schist and not the Chapolapa Formation.

The Taxco Viejo Greenstone unconformably covers the Taxco Schist. This relation is either the expression of a flat tectonic contact between the two, or it reflects that the Taxco Viejo Greenstone is the youngest lithologic unit within the Taxco Schist. The Taxco Viejo Greenstone was sampled at three localities (Figure 1, localities B, C, and D), which are

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Figure 1.- Map showing the localities where volcanic rock samples (black triangles with capital letters) were collected from the younger greenschist-facies package of the pre-Upper Jurassic and/or Lower Cretaceous basement in south-central Mexico.

Table 1.- Major- and trace-element analytical data of six volcanic rock samples from the younger, greenschist-facies package of the pre-Upper Jurassic and/or Lower Cretaceous basement in south-central Mexico. For location of samples see text and Figure 1.

SAMPLE	Α	в	С	D	E	F
SiO ₂ [%]	58.24	49.80	59.42	61.72	61.41	54.64
TiO ₂	0.76	1.18	0.90	0.97	0.79	0.96
Al ₂ O ₃	16.18	14.85	16.78	15.23	17.00	16.24
Fe ₂ O ₃	3.66	1.60	2.01	2.53	5.84	2.32
FeO	3.96	8.04	2.66	1.69	0.45	5.38
MnO	0.14	0.16	0.06	0.07	0.09	0.14
MgO	3.88	6.55	4.84	3.00	2.77	5.45
CaO	6.85	8.39	3.11	6.45	1.04	7.49
Na ₂ O	2.85	2.68	3.71	3.08	6.11	4.38
K ₂ O	1.50	1.63	2.18	1.45	2.58	0.48
P ₂ O ₅	0.08	0.20	0.14	0.17	0.14	0.14
LOI	2.67	4.18	3.34	3.69	2.28	2.47
Total	100.61	98.86	98.87	99.71	100.22	99.79
Ba [ppm]	681	407	585	597	685	283
Nb	17	24	8	17	12	10
Rb	32	50	39	29	82	11
Sr	365	269	339	513	209	268
Y	22	20	17	18	37	23
Zr	159	136	177	150	258	146
La	10	15	11	10	18	9
Ce	19	30	23	21	38	17
Nd	<10	<10	<10	12	<10	11
Sm	3.3	3.9	3.2	3.3	5.3	3.1
Eu	0.9	1.2	1.0	1.0	0.8	0.9
Yb	2	2	2	2	3	2
Lu	0.3	0.2	0.2	0.2	0.4	0.3
Sc	21.6	28.3	19.6	22.4	20.4	30.5
Tb	<1	<1	<1	<1	<1	<1
Th	2.5	3.2	3.0	2.3	8.5	2.1
Tm	<10	<10	<10	<10	<10	<10
U	1	1	1	<1	2	<1
Cs	1.5	4.9	2.0	1.4	4.2	<0.5
Hf	2.5	2.0	3.1	2.4	4.6	2.0
Та	<1	<1	<1	1.1	<1	<1

LOI: Loss on ignition at 950°C.

Major-element analyses were performed by A. Lozano-Cobo in the Geochemistry Laboratory of the Instituto de Geología, Universidad Nacional Autónoma de México.

Trace-element analyses were performed in the North Vancouver, Canada, laboratories of Bondar Clegg & Co.

within the confinement of the Taxco quadrangle (de Cserna and Fries, 1981). The first of these is about one kilometer north of Zacualpan along the road to the El Regenerador adit (Díaz-García, 1980); the sample was obtained from a poorly foliated andesite metalava (Table 1, sample B). The second and third localities are practically side by side, about 3 km east of Tonatico along the highway to Taxco. In this area, the outcropping lithology of the Taxco Viejo Greenstone consists of foliated metalahar. A representative sample of the matrix of this lahar (Table 1, sample C) resulted in an andesitic to dacitic composition, while a representative clast, not as well foliated as the matrix, turned out to be a dacite (Table 1, sample D).



Figure 2.- AFM diagram showing typical calc-alkaline differentiation trend of volcanic rocks in the younger greenschist-facies package of the pre-Upper Jurassic and/or Lower Cretaceous basement in south-central Mexico. The samples identified by capital letters are the same as those in Figures 1 and 3, and in Table 1.

The Chapolapa Formation is the so far known youngest lithostratigraphic unit of the younger, greenschist-facies package of the pre-Upper Jurassic and/or Lower Cretaceous basement in south-central Mexico. At places, the lower parts of the Chapolapa Formation appear to change laterally into the Taxco Viejo Greenstone (de Cserna, 1965); however, much more work is needed to prove such relations conclusively. A sequence of redbeds with interbedded lava flows of the Chapolapa Formation has excellent exposures along the bed of river Chapolapa in the outskirts of Acahuizotla (de Cserna, 1965; Figure 1, locality E), about 25 km SSE of Chilpancingo in south-central State of Guerrero. At this locality, the sampled volcanic rock (Table 1, sample E) is from a dark gray dacite flow that is interbedded in a sequence of reddish-purple, poorly sorted conglomerate whose pebbles are poorly rounded milky quartz; the entire sequence has a weakly developed foliation.



Figure 3.- MORB-normalized trace element variation diagram for volcanic rock samples from the younger greenschist-facies package of the pre-Upper Jurassic and/or Lower Cretaceous basement in south-central Mexico. Samples identified by capital letters are the same in Figures 1 and 2, and in Table 1.

CONCLUSIONS

The analytical results obtained on these rock samples are given in Table 1. The major element analytical data were plotted on the classical AFM diagram (Figure 2) on which the typical calc-alkaline trend can be observed. The trace element analytical data were plotted on a MORB-normalized trace element variation diagram (Figure 3). It can be observed, on this diagram, the significant enrichment in Sr, K, Rb, Ba, and Th, which, because of their low ionic potential, tend to mobilize by aqueous solutions in compressive environments associated with convergent plate margins (Pearce, 1982).

Considering the above new data, the authors feel that the earlier proposed continental growth for south-central Mexico through the welding of successive orogenic belts from the west still appears to be a feasible geotectonic model (de Cserna, 1958, 1965, 1976; Rogers *et al.*, 1974), and that the more recently proposed models continue being hypothetical until greater extensions are covered with good quality geological mapping and the supporting geochemical investigations.

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